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Nitrification and denitrification response to varying periods of desiccation and inundation in a western Kansas stream

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Abstract Changing environmental conditions and increased water consumption have transformed many historically perennial stream systems into intermittent systems. Multiple drying and wetting events throughout the year might impact many stream processes including nitrification and denitrification, key components of the nitrogen (N) cycle. During summer 2007, an experimental stream was used to dry and then rewet stream sediments to determine the effects of desiccation and rewetting of stream sediment on nitrification and denitrification potentials. Mean (±SE) nitrification and denitrification rates in sediment not dried (controls) were $0.431 \pm 0.017 \ \mu g$ NO_3^- -N/cm²/h and 0.016 ± 0.002 µg N₂O-N/cm²/h, respectively. As sediment samples dried, nitrification rates decreased. Rates in sediments dried less than

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Department of Biological Sciences, University of Arkansas-Fayetteville, 601 Science and Engineering, Fayetteville, AR 72701, USA e-mail: bjaustin@uark.edu 7 d recovered to levels equal or greater than those in the controls within 1 d of being rewetted. Denitrification rates were not affected by 1 d of drying, but samples dried greater than 1 d experienced reduced rates of denitrification. Denitrification in sediments dried 7 d or less recovered by day seven of being rewetted. Nitrification and denitrification processes failed to fully recover in sediments dried more than 7 d. These results demonstrate that alterations in stream's hydrology can significantly affect N-cycle processes.

Keywords Nitrification · Denitrification · Desiccation

Introduction

Intermittent streams are common in grassland and arid biomes because of prevailing weather patterns and land/water uses (Dodds et al., 2004). This pattern is certainly relevant in the Great Plains of central North America where drought conditions over the past several decades, coupled with increased groundwater use, have resulted in lowered water tables across the region (Cross & Moss, 1987; Angelo, 1994; Eberle et al., 1998; Gido et al., 2010). This increase in groundwater usage has decreased discharge rates and has dramatically reduced the number of perennial streams in western Kansas (Angelo, 1994; Sophocleuous, 2000; Dodds et al., 2004; Gido et al., 2010). There is speculation that this drying trend may continue and perhaps expand because drought conditions have been predicted to worsen through the twenty-first century resulting in a 10–30% decrease in runoff in mid-latitude and dry tropical areas worldwide due to global climate change (IPCC, 2007). Aquatic ecologists have long studied the ecology of intermittent streams and the organism inhabitants (primarily macroinvertebrates and fish), but much less in known about the effects of periodic drying on microbial processes, e.g., cycling of nitrogen (N) in these systems.

Biologically available N enters landscapes via several pathways such as microbial N fixation, atmospheric deposition, agricultural inputs, and municipal and industrial waste. Excess N within an ecosystem that is not assimilated is susceptible to export through runoff into streams and rivers or through leaching into groundwater. Recent research has shown that small streams can be major areas for nutrient cycling, decreasing export of N, which potentially decreases eutrophication and hypoxia downstream (Alexander et al., 2000; Peterson et al., 2002; Mulholland et al., 2008). Several inherent characteristics of smaller systems predispose them to relatively high N processing rates. For example, small streams generally have complex channels, shallow water, and long retention times, which increase contact between dissolved nutrients and stream substrates (Bernot & Dodds, 2005). Although most N-cycling processes occur in small streams, the predominant processes which ultimately result in the conversion of N from a biologically to nonbiologically available form are nitrification and denitrification.

Nitrification, the chemoautotrophic oxidation of ammonium (NH₄⁺), is an aerobic two-step process conducted by relatively few bacterial species that belong to two different functional groups: NH₄⁺ oxidizers, which oxidize NH₄⁺ to nitrite (NO₂⁻) and NO₂⁻ oxidizers, which oxidize NO₂⁻ to nitrate (NO₃⁻) (Bothe et al., 2000; Konhauser, 2007). Denitrification is the anaerobic reduction of NO₃⁻, which is produced via nitrification, and other oxidized N species to dinitrogen gas (N₂). Unlike nitrification, there are multiple orders of bacteria capable of conducting denitrification include NH₄⁺ and dissolved oxygen (DO) availability (Triska et al.,

1990; Kemp & Dodds, 2001), and organic carbon quality and availability (Strauss & Dodds, 1997; Strauss & Lamberti, 2000). While regulating factors of denitrification include NO_3^- and DO availability (Nielsen et al., 1990), organic carbon availability (Hedin et al., 1998), and temperature (Andersen, 1977). Changes in water content of substrates and its impacts on nitrification and denitrification have also been examined.

Several studies have also focused on the effects of drying and rewetting on nitrification and denitrification among several different habitats and substrate types. Smith & Parsons (1985) looked at the response of denitrification to desiccated soils, finding increased denitrification rates following re-moistening. Mitchell & Baldwin (1999) measured nitrification and denitrification in desiccated lake sediments, finding nitrification to be non-detectable: however, denitrification was not impacted. While Qiu & McComb (1996) found increased NH_4^+ availability due to mineralization during the drying process and enhanced nitrification rates upon rewetting of dried sediment cores collected from a wetland, due aerobic conditions. Only one study has focused on the desiccation of a lotic system and the impacts on nitrification and denitrification. Cavanaugh et al. (2006) measured nitrification and denitrification rates during and after the intentional water-level draw down in the Upper Mississippi River (UMR), finding decreased nitrification rates due to competition between nitrifiers and macrophytes for NH4⁺ and decreased denitrification due to increased aerobic conditions.

Other studies have found a decrease in bacterial biomass after sediments dried and a subsequent rise in NH_4^+ availability and nitrification upon rewetting (Groffman & Tiedje, 1988; Qiu & McComb, 1996; Mitchell & Baldwin, 1998). The resulting NO_3^- could, in turn, promote denitrification in nearby anaerobic areas (Kern et al., 1996). The findings of both increased nitrification and denitrification might suggest significant reductions in TN for certain systems upon rewetting.

The objectives of this study were to determine how periodic drying of stream sediments affects nitrification and denitrification, and determine how quickly nitrification and denitrification rates return to preperturbation levels after rewetting. We hypothesized that nitrification rates would increase in response aerobic conditions during drying and that $NO_3^$ would accumulate, thereby stimulating denitrification when sediments rewetted and anaerobic conditions returned. We also predicted that nitrification and denitrification recovery time, after sediments were rewetted, would depend on the time desiccated, based on the assumption that the impact on the microbial communities would be directly proportional to the time dried.

Materials and methods

Site description

This study was conducted at a Fort Hays State University cattle research site on the Saline River in northeastern Ellis County, Kansas (latitude: 39°4′4.97″N; longitude: 99°6′13.73″W). The Saline River falls within the Smoky Hill—Saline River basin and is one of the five major tributaries of the Kansas River. The general land use upstream of the site is mostly ungrazed and grazed rangeland. The Saline River maintains a low base flow (0.10 m³/s) throughout most of the year; however, during periods of high precipitation, discharge increases rapidly.

Project description

The experimental stream consisted of a fiberglass lined basin $(L \times W \times D = 274 \text{ cm} \times 61 \text{ cm} \times 18 \text{ cm})$. Stream water from the Saline River was supplied to the experimental stream at a rate of 0.75–1.8 l/min via a battery powered peristaltic pump (Master-Flex L/S model # 77200-62, Barnant Co., USA). Stream water was distributed evenly over the experimental stream with a 1.27 cm diameter PVC pipe with holes drilled along its length. The drying table was positioned adjacent to the experimental stream and was constructed to similar dimensions as the experimental stream, with the only variation being holes were drilled into the bottom to facilitate drainage of drying sediments and to drain rain water.

Sediment used in this study was collected from the Saline River adjacent to the experimental stream, and was predominantly a mixture of sand and silt. Approximately 0.065 m^3 of sediment was collected from the stream (only sediment from the top 5 cm

was collected), placed into a large container and homogenized before randomly distributing into 252 sample cups ($\sim 250 \text{ cm}^3$). Sample cups consisted of polycarbonate rings (7.6 cm diameter \times 7 cm tall) fitted with 0.5 mm fiberglass mesh on the bottom that limited sediment leakage, but allowed water to drain through the sediment and out of the cups. The sample cups were then placed into the experimental stream for 1 week to acclimate.

After the acclimation period, the 252 sample cups were randomly distributed as follows. Seven sample cups were taken to the laboratory to be analyzed as the initial control, 35 sample cups were left in the experimental stream to act as wet controls, and 210 sample cups were moved to the drying table and allowed to drain dry. The sediment samples placed on the drying table were randomly divided into treatments based on the number of days they were to be dried (1, 3, 7, 14, 21, and 28 d), and within each of those drying treatments they were further divided into treatments based on the number of days they were to be rewetted (0, 1, 3, 7, 14, 21, and 28 d) with five replicates for each combination of treatments.

At each designated treatment interval, five replicates (sample cups) were collected and returned to the laboratory for sediment and N-cycling analyses. Processing of all samples was initiated within 6 h of collection. Sediment properties measured included: percent water, ash free dry mass (AFDM), exchangeable NH_4^+ concentration, and rates of nitrification and denitrification. Water samples were also collected daily from the Saline River and the experimental stream to compare concentrations of NH_4^+ , $NO_2^- + NO_3^-$, and soluble reactive phosphorus (SRP) between the actual and experimental streams. Temperature, pH, conductivity, DO, and DO percent saturation were monitored at 15-min intervals throughout the study using Manta multiparameter probes (Eureka Environmental Inc., Austin, TX) placed in both the experimental stream and the Saline River. Multiparameter probe data and water for laboratory analyses were collected in the stream at the point of intake for the experimental stream. Immediately upon arrival to the laboratory, water samples (100 ml) were filtered through dried and pre-weighed filters (Whatman GF/F), preserved with 50 μ l of 1 N H₂SO₄, and refrigerated until colorimetric analyses of NH_4^+ , $\mathrm{NO_2}^- + \mathrm{NO_3}^-,$ and SRP were conducted. The filter was used for determining total suspended solids (TSS).

TSS in the filtered volume were quantified by drying the filter at 50° C for 24 h and then weighing.

Sediment and water characteristics

Sediment moisture was determined on a 20 cm³ subsample of the sediment returned to the laboratory by the mass difference measured between wet and dried samples (24 h at 105°C). AFDM was determined by subtracting ash mass (3 h at 500°C) from dry mass. Sediment moisture content and bulk density were calculated following Håkanson & Jansson (1983). Exchangeable NH_4^+ was determined in conjunction with nitrification by measuring NH_4^+ in a potassium chloride (KCl) extract via the phenate method (Solorzano, 1969). Surface water SRP, $NO_2^- + NO_3^-$, and NH_4^+ concentrations were determined by using the standard colorimetric molybdate/ascorbic acid, cadmium reduction, and phenate methods, respectively (APHA, 2005), and using Turner SP-890 UV-VIS spectrophotometer а (Barnstead International, Dubuque, IA).

Nitrification

Nitrification rates were determined by using a modification of the nitrapyrin inhibition method (as described in Strauss et al., 2004). Two slurries for each sediment sample [25 cm³ sediment and 81 ml filtered surface water or DI water (dry samples)] were incubated aerobically in the dark (3 d) on an orbital shaker table (175 rpm). In each pair of slurries, nitrification was inhibited in one flask by adding 20 µl nitrapyrin/dimethyl-sulfoxide (DMSO) solution (final nitrapyrin concentration in flask = 10 mg/l). In the control flask, 20 µl DMSO without nitrapyrin was added. KCl extraction of NH4⁺ was conducted on a 6 ml sample of each slurry before and after the 3 d incubation period. Ammonium extracts were stored at 4°C until analysis. Ammonium concentration (phenate method) of KCl extracts was measured within 1 week of sample collection and used to calculate gross nitrification rates (Strauss & Lamberti, 2000). Nitrification rates were determined to be the difference between the change of NH_4^+ in the nitrapyrin addition and the change of NH_4^+ in the DMSO control (nitrification rate = $([N_{final} - N_{initial}])$ - [D_{final} - D_{initial}]/incubation period). Where N is NH_4^+ concentration in the flask containing nitrapyrin addition and D is the NH_4^+ concentration in the flask containing the DMSO control. The NH_4^+ analysis of the initial samples containing only DMSO was also used as the measure of exchangeable NH_4^+ for that particular sediment sample.

Denitrification

Denitrification in the sediment slurries was determined by using the denitrification enzyme activity (DEA) metric (Groffman et al., 1999; Richardson et al., 2004; Strauss et al., 2006), which is a variation of the commonly used acetylene (C_2H_2) block technique (Sorensen, 1978; Richardson et al., 2004). DEA determines the potential for microorganisms to reduce NO_3^- to N_2 under optimal conditions, and is an indicator for measuring differences in denitrification potential between treatments (Groffman et al., 1999).

Denitrification enzyme activity slurries contained 20 cm³ sediment, 20 ml unfiltered surface water (or deionized [DI] water for dry samples), and 5 ml DEA solution (final slurry concentrations = 100 mg/l chloramphenicol, 12 mg/l glucose-C, and 14 mg/l potassium nitrate [KNO₃–N]). The antibiotic chloramphenicol inhibits production of additional denitrifying enzymes without killing bacteria or destroying enzymes already present. Glucose provides an energy source and KNO₃ an electron acceptor suitable for denitrification.

Slurries were placed in 237 ml glass canning jars and sealed with standard canning rings and lids equipped with 13 mm butyl septa for gas additions and removal. Jars were evacuated for 5 min using a vacuum pump, and the negative pressure was equalized with chromatographic grade helium (He) gas (99.999% pure) to initiate anaerobic conditions. Atomic absorption grade C_2H_2 (40 cm³) was then added to the headspace of sealed incubation jars to inhibit nitrous oxide (N2O) reduction to N2 gas. Slurries were incubated at room temperatures in the dark on an orbital shaker table (175 rpm) (New Brunswick Scientific C10 Platform Orbital Shaker). At specific intervals (10, 30, 60, and 90 min from the time of C₂H₂ addition) headspace gas samples (10 ml) were extracted from the jars and placed into separate evacuated 10 ml glass serum bottles for later analysis of N₂O concentration. Nitrous oxide concentrations were measured within 1 week of collection on an SRI 8610C gas chromatograph (SRI Instruments, Inc., Torrance, CA) equipped with a 63 Ni electron capture detector, and argon-5% methane (P5) was used as the carrier gas. The standard curve used to determine sample concentrations was produced using Excaliber 0.95, 10, 100 ppm N₂O gas standards (Linweld Inc., Lincoln, NE). DEA rates were calculated from the linear increase in N₂O concentration throughout the incubation period following the equations in Groffman et al. (1999). Volume/volume concentrations and adjusted to incorporate the volume of N₂O dissolved in liquid phase at the given pressure and temperature using the ideal gas law and the Bunsen solubility coefficient, respectively.

Data analysis

All statistical analyses were conducted by using SAS statistical software (version 9.1, SAS Institute, Cary, NC). The Shapiro-Wilk test and visual inspection of normal probability plots were used to assess normality of the data sets prior to running parametric tests to look at the relationships among these data. After determining that data sets were distributed normally, a parametric general linear model one-way analysis of variance (ANOVA; F-test) was used to identify significant treatment effects. If the overall F-test was significant, the LS means procedure was used to determine specific differences among treatment means. In addition, Pearson's correlation coefficients were calculated to explore relationships between nitrification and DEA rates with sediment moisture (% H_2O), exchangeable NH₄⁺, and AFDM. All comparisons and correlations were evaluated at $\alpha = 0.05$. The main purpose of the statistical analyses was to determine the effects of desiccation on the major processes (nitrification and denitrification), and then the subsequent recovery of major processes once sediments were re-inundated. An a priori definition of recovery was used, where treatment rates return to or exceed rates measured in the controls.

Results

Experimental stream versus stream

Experimental stream and stream water nutrient chemistry (SRP, NH_4^+ , $NO_2^- + NO_3^-$) were similar. However, TSS in the Saline River was consistently higher than the experimental stream (Fig. 1). Multiparameter probe data for all parameters (conductivity, temperature, pH, and DO) exhibited higher values for conductivity and greater diel shifts for temperature, pH, and DO in the experimental stream compared to the Saline River (Fig. 2).

Effect of drying on sediment characteristics

In addition to %H₂O, drying of the sediments also affected exchangeable NH₄⁺ concentration (Fig. 3). Percent H₂O decreased to below 5% by 3 d of drying and remained below 5% for the rest of the drying period. In contrast, exchangeable NH₄⁺ availability increased over the 28 d drying period, increasing from control sediments (0.112 \pm 0.003 µg NH₄⁺–N/g dw) to sediments dried 28 d (0.565 \pm 0.037 µg NH₄⁺–N/g dw). AFDM was variable over the 28 d drying period ranging from 0.020 to 0.038 mg AFDM/cm³.

Effect of rewetting on sediment characteristics

We examined sediments dried for 28 d prior to rewetting to assess the effect of rewetting on sediment characteristics. As sediments rehydrated, the response was opposite to that of drying for exchangeable NH_4^+ ; exchangeable NH_4^+ decreased slowly to 0.266 \pm 0.023 µg NH_4^+ /g dw by the 28th d of rewetting (Fig. 4). There was no definitive correlation between number of days rewet and AFDM.

Effect of drying on nitrification and denitrification

Drying of sediment samples affected rates of nitrification and DEA similarly. For nitrification, as sediments dried there was a significant decline from control rates ($0.430 \pm 0.017 \ \mu g \ N/cm^2/h$) to rates found in sediments dried 1 d ($0.108 \pm 0.043 \ \mu g \ N/cm^2/h^1$) ($F_{(1,10)} = 61.57$; P < 0.0026). There were no significant differences found among any of the drying intervals from 1 to 28 d. One day of drying did not change DEA rates from those of the control ($0.016 \pm 0.002 \ \mu g \ N/cm^2/h$), but DEA rates declined sharply to non-detectable levels after 3 d of drying. DEA rates between 3 and 28 d of drying did not differ significantly. Fig. 1 Total suspended solids (TSS; A), and concentrations of $NH_4^+(B)$, $NO_3^- + NO_2^-(C)$, and soluble reactive phosphorus (SRP; D) measured in the Saline River (*open circle*) and from the experimental stream (*filled circle*) used in this study

Fig. 2 Temperature (A), specific conductivity (B), dissolved oxygen (C), and pH (D) measured in the Saline River (*open circle*) and from the experimental stream (*filled circle*) used in this study. Plots represent a 1 week subset from 11 to 18 August 2007



Effect of rewetting on nitrification and denitrification

Inundation of sediment samples produced varying results between the two processes. Nitrification recovered in sediments dried for 1, 3, 7, and 21 d. However, after rates in each group recovered, the rates declined back below control rates. Rates in

sediments dried 14 and 28 d increased toward recovery, but did not reach the initial rate set by the control by the end of the 28 d of rewetting (Fig. 5).

Denitrification enzyme activity rates from rewetting of samples dried for only 1 d were not significantly different than the control with the exception of 21 d rewet which was higher. The sediments dried for 3, 7, and 21 d recovered by 7 d of being rewetted.



Fig. 3 Sediment exchangeable NH_4^+ and $\%H_2O$ measured within Saline River sediments as they dried. The control indicates sediments that were not dried. *Error bars* = ±1 SE



Fig. 4 Sediment exchangeable NH_4^+ and $\%H_2O$ measured in Saline River sediments dried 28 d during rewetting process. *Error bars* = ± 1 SE

DEA samples dried 14 and 28 d failed to recover by the 28th d of being rewetted (Fig. 6).

Correlations

As sediment samples dried, both nitrification and DEA rates declined. During this time both nitrification and DEA were positively correlated with %H₂O (r = 0.632,P < 0.001; r = 0.888,P < 0.001, respectively). DEA also was negatively correlated with exchangeable NH_4^+ (*r* = -0.684, *P* < 0.001) and positively correlated with AFDM (r = 0.506, P = 0.002). During the rewetting process nitrification was negatively correlated with DEA (r =-0.204, P = 0.006) and DEA was negatively correlated with exchangeable NH_4^+ (r = -0.512, P < 0.001) (Table 1). No other significant correlations were observed for either process during the rewetting period.

Discussion

Alterations in the physical environment, including water availability, can result in profound changes in the abundance, diversity, and activity of the microbial community (Qiu & McComb, 1996; Mitchell & Baldwin, 1998). Several studies have examined this interaction between hydrological fluctuations on one or both of these N-cycling processes, in habitats including soil (Groffman & Tiedje, 1988), lake sediment (Qiu & McComb, 1996; Mitchell & Baldwin, 1999), floodplains (Kern et al., 1996) and large order streams and backwater areas (Cavanaugh et al., 2006). Although small headwater streams are major areas of nutrient cycling (Alexander et al., 2000; Peterson et al., 2002; Bernot & Dodds, 2005; Mulholland et al., 2008), to our knowledge, no studies examining the impact of desiccation of headwater streams on nitrification and denitrification have been conducted.

In comparison to rates measured in other studies that have used similar methods, our measured mean rate of nitrification (0.430 µg N/cm²/h for wet controls) was similar to rates measured in two intermittent tallgrass prairie streams $(0.22-0.46 \ \mu g \ N/cm^2/h)$ (Kemp & Dodds, 2002), three agricultural streams from various regions of the United States $(0.16-0.44 \ \mu g \ N/cm^2/h)$ (Duff et al., 2008), and the sandy main channel of the UMR (0.35 μ g N/cm²/h) (Strauss et al., 2004). However, our rate was much higher than those published by Strauss et al. (2002) for streams in northern Wisconsin and the upper peninsula of Michigan (0.015–0.089 µg N/cm²/h) and Starry et al. (2005) for a small forested headwater stream in North Carolina (0.028–0.081 µg N/cm²/h). Our DEA control rate (0.016 µg N/cm²/h) was relatively low compared to rates in similar studies. DEA rates of 0.41–2.02 μ g N/cm²/h were reported for the three agricultural streams in the study by Duff et al. (2008) and 0.178–22.2 μ g N/cm²/h were reported for the UMR (Richardson et al., 2004). The relative magnitude of the process rates we measured was likely a reflection of the availability of nutrients in the stream. The low DEA rates can be explained by the low stream water NO3⁻ concentration. Whereas, the higher nitrification rate was a result of the higher stream water NH₄⁺ concentration complimenting the NH_4^+ available in the sediment.

During the sediment drying and rewetting process, there are several key changes that occur that might Fig. 5 Nitrification rates measured in Saline River sediments during the rewetting process. The number of days dried is shown in separate panels: 1 (A), 3 (B), 7 (C), 14 (D), 21 (E), and 28 d (F). F-test results shown are for the overall one-way ANOVA model. Different letters associated with the treatment means (bars) signify significant difference (P < 0.05; LS means procedure) among the means. Error $bars = \pm 1$ SE



affect nitrification and DEA potentials. Studies examining dried soil and sediments that were rewetted found increased organic carbon (Kieft et al., 1987; Groffman & Tiedje, 1988) and $\rm NH_4^+$ (Groffman & Tiedje, 1988; Qiu & McComb, 1996) concentrations upon rewetting, likely due to osmotic lysis of bacterial cells (Kieft et al., 1987; Groffman & Tiedje, 1988; Qiu & McComb, 1994, 1995; Mitchell & Baldwin, 1998; Modini et al., 2002). Although there were no direct measures of bacterial biomass in this study, an increase in exchangeable NH_4^+ was observed as sediments were dried (Fig. 3) as seen by others. It is likely that the increase in exchangeable NH_4^+ was due to the osmotic lysis of bacterial cells

Fig. 6 DEA rates measured in Saline River sediments during the rewetting process. The number of days dried is shown in the separate panels: 1 (A), 3 (B), 7 (C), 14 (**D**), 21 (**E**), and 28 d (**F**). F-test results shown are for the overall one-way ANOVA model. Different letters associated with the treatment means (bars) signify significant difference (P < 0.05; LS means procedure) among the means. Error $bars = \pm 1$ SE



resulting in decreased bacterial biomass, which could possibly affect N processing rates. However, previous studies have found substrate nutrient concentrations and habitat conditions (temperature, O_2 , and redox potential) to be more important with respect to microbial processing rates (García-Ruiz et al., 1998; Pattinson et al., 1998; Wall et al., 2005) than bacterial biomass (Jones, 1979; Davidson et al., 1985; Hermansson et al., 2004; Rigobelo & Nahas, 2004). The drying of sediments likely increases O_2 permeability into the sediment which should favor nitrification over DEA during the initial desiccation period (Mitchell & Baldwin, 1999), which is what we had predicted to occur. However, we found %H₂O to be the independent variable regulating nitrification and DEA rates during the initial days of drying. After the first day roughly 15% H₂O remained in the sediment cores and at this point nitrification rates

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Table 1 Pearson correlations table for nitrification, DEA, and the sediment properties measured in the experimental stream sediments during the drying and rewetting process		Nitrification	DEA	%H ₂ O	AFDM	Exc. NH ₄ ⁺
	All data					
	Nitrification	1.0	-0.1164	0.1632	0.1117	-0.1232
		0	0.0870	0.0166	0.1047	0.0700
	DEA	_	1.0	0.3890	0.07635	-0.5165
		_	0	< 0.0001	0.2684	< 0.0001
	%H ₂ O	_	_	1.0	0.1200	-0.20643
		_	_	0	0.0829	0.0023
	AFDM	_	_	_	1.0	0.0890
		_	_	_	0	0.1967
	Wet data					
	Nitrification	1.0	-0.2042	0.0676	0.0978	-0.1240
		0	0.0060	0.3689	0.1965	0.0973
	DEA	_	1.0	0.1151	-0.0038	-0.5120
		_	0	0.1249	0.9603	< 0.0001
	%H ₂ O	_	_	1.0	-0.0814	-0.1955
		_	_	0	0.2845	0.0087
	AFDM	_	_	_	1.0	0.1754
		_	_	_	0	0.0199
	Dry data					
	Nitrification	1.0	0.5628	0.6323	0.1852	-0.2440
		0	0.0003	< 0.0001	0.2796	0.1455
	DEA	_	1.0	0.8879	0.5062	-0.6842
		_	0	< 0.0001	0.0016	< 0.0001
For each variable pair, the upper number is the Pearson correlation coefficient (<i>r</i>) and the lower number is the <i>P</i> -value associated with <i>r</i>	%H ₂ O	_	_	1.0	0.4748	-0.6165
		_	_	0	0.0040	< 0.0001
	AFDM	_	-	-	1.0	-0.4188
		_	_	-	0	0.0110

with r decreased. It was not until the third day of drying and sediment core %H₂O dropping below 5% that DEA rates declined. This suggests that the denitrifiers present in this system are more diverse and or drought tolerant than the nitrifiers. Although %H₂O appears to be the regulating factor of N-cycle transformations during drying, O₂ availability was

more important during the rewetting process.

Recovery of nitrification rates in other studies has been variable. Increased rates of nitrification were found in lake sediments following re-flooding (Qiu & McComb, 1996) and upon rewetting of dried forest soils (Fierer & Schimel, 2002). Nitrification potential, however, was found to decrease following the end of the intentional drawdown of the UMR (Cavanaugh et al., 2006) and Zaman & Chang (2004) found a 30+ d lag between rewetting and rates returning to normal.

As predicted, we found that sediments dried for a short period of time, 7 d or less, nitrification rates recovered within 1 d of rewetting, followed by a decline to pre-recovery rates (Fig. 5A-C). This pattern is most likely the result of a switch from an aerobic to anaerobic environment in the sediment cores. Although the oxygen content of the cores was not measured at any point throughout the study, findings from DEA measurements supports this hypothesis. Sediments dried for greater than 7 d, although having increased NH4+ availability, were unable to recover by 28 d of rewetting. There was one outlier in the group of treatments dried greater than 7 d, the treatment 21 d dry, fully recovered by 3 d of rewetting; this outlier will be discussed in more detail later.

As seen with nitrification rates, there was a lag between rewetting and recovery for DEA rates as well. Sediments dried 3 and 7 d, recovered within 7 d of rewetting. However, it was not until nitrification rates recovered and then began declining that DEA rates recovered, in the 3 and 7 d dry treatments. In addition, sediments dried only 1 d showed an increase in rates following the decline of nitrification rates. This suggests as nitrification commenced following rewetting of the sediment cores, NH_4^+ and O_2 in the cores were used, producing NO₃⁻ and an anaerobic environment, creating the appropriate conditions for denitrification to occur. These conditions would have promoted the production of denitrifying enzymes which can persist even after all the NO_3^- is used. The persistence of these enzymes and the NO₃⁻ source in the DEA solution allowed measured DEA rates to remain relatively stable for several days following recovery. As with nitrification rates, DEA rates also recovered in the 21 d dry treatment only after nitrification rates began to decline, and failed to fully recover in the 14 and 28 d dry treatments.

The recovery of nitrification and DEA rates in sediments dried 21 d was unexpected since both processes failed to recover in sediments dried 14 d. One possible explanation for this is that the 14 d dry treatment could have recovered and then declined sometime between 1 and 3 d of being rewetted and was missed due to the sampling frequency. Evidence for this occurring can be seen in Fig. 6D. Although DEA rates never fully recovered in sediments dried 14 d, the increase in DEA rates seen 3-7 d after rewetting would not have occurred unless nitrification was also occurring. If the recovery of sediments dried 14 d was not detected as suggested, then the 21 d dry treatment would not be an outlier. The same logic could be used to suggest that perhaps some level of nitrification recovery might have also occurred in the 28 d dry treatment.

Sediment drying can cause a loss of microbial biomass and subsequent mineralization of N and C, resulting in increased levels of NH_4^+ in sediments. It was predicted that this increase in NH_4^+ would promote elevated nitrification rates upon re-inundation of sediments, increasing the availability of NO_3^- to denitrifiers. Ammonium did increase in the sediments due to desiccation (Fig. 3); however, much of this was probably lost to the water column upon rewetting, resulting in very little NH_4^+ being converted all the way through to N_2 (Fig. 1B). This increase in NH_4^+ in the water column could

potentially be transported downstream to a larger tributary where it would have a lower probability of being processed.

This pulse of NH_4^+ downstream would only be an issue following re-inundation of the stream channel. While the flow of streams in the future is uncertain, some climatic studies suggest the arid conditions of many regions throughout the world will continue or worsen due to decreased precipitation and increased evaporation (National Assessment Synthesis Team, 2001; IPCC, 2007). Clearly, if arid conditions predominate in the future, altered hydrological conditions of the region will not be the only effect. Multiple ecosystem processes, not just nitrification and denitrification, will be severely altered.

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References

- Alexander, R. B., R. A. Smith & G. E. Schwarz, 2000. Effect of stream channel size on the delivery of nitrogen to the Gulf of Mexico. Nature 403: 758–761.
- Andersen, J. M., 1977. Rates of denitrification of undisturbed sediments of six lakes as a function of nitrate concentration, oxygen, and temperature. Archive fur Hydrobiologia 80: 147–159.
- Angelo, R. T., 1994. Impacts of declining streamflow on surface water quality. 11th Annual Water and the Future of Kansas Conference Proceedings, Manhattan, KS: 1–2.
- APHA (American Public Health Association), 2005. Standard Methods for the Examination of Water and Wastewater, 21st ed. American Public Health Association, Washington, DC: 1368 pp.
- Bernot, M. J. & W. K. Dodds, 2005. Nitrogen retention, removal, and saturation in lotic ecosystems. Ecosystems 8: 442–453.
- Bothe, H., G. Jost, M. Schloter, B. B. Ward & K. P. Witzel, 2000. Molecular analysis of ammonia oxidation and denitrification in natural environments. FEMS Microbiology Reviews 24: 673–690.
- Cavanaugh, J. C., W. B. Richardson, E. A. Strauss & L. A. Bartsch, 2006. Nitrogen dynamics in sediment during water level manipulation on the Upper Mississippi River. River Research Applications 22: 651–666.
- Cross, F. B. & R. E. Moss, 1987. Historic changes in fish communities and aquatic habitats in plains streams of Kansas. In

Matthews, W. J. & D. C. Heins (eds), Community Evolutionary Ecology of North American Stream Fishes. University of Oklahoma Press, Norman, OK: 155–165.

- Davidson, E. A., M. K. Strand & L. F. Galloway, 1985. Evaluation of the most probable number method for enumerating denitrifying bacteria. Soil Science Society of American Journal 49: 642–645.
- Dodds, W. K., K. Gido, M. R. Whiles, K. M. Fritz & W. J. Matthews, 2004. Life on the edge: the ecology of Great Plains prairie streams. BioScience 54: 205–216.
- Duff, J. H., A. J. Tesoriero, W. B. Richardson, E. A. Strauss & M. D. Munn, 2008. Whole-stream response to nitrate loading in three streams draining agricultural landscapes. Journal of Environmental Quality 37: 1133–1144.
- Eberle, M. E., T. L. Wenke & N. E. Mandrak, 1998. Assessment of the Solomon River, North Fork Solomon River, and South Fork Solomon River in Northwestern Kansas. Kansas Department of Wildlife and Parks and U.S. Bureau of Reclamation: 63.
- Fierer, N. & J. P. Schimel, 2002. Effects of drying-rewetting frequency on soil carbon and nitrogen transformations. Soil Biology and Biochemistry 34: 777–787.
- García-Ruiz, R., S. N. Pattinson & B. A. Whitton, 1998. Denitrification in river sediments: relationships between process rates and properties of water and sediment. Freshwater Biology 39: 467–476.
- Gido, K. B., W. K. Dodds & M. E. Eberle, 2010. Retrospective analysis of fish community change during a half-century of land-use and streamflow changes. Journal of the North American Benthological Society 29: 970–987.
- Groffman, P. M. & J. M. Tiedje, 1988. Denitrification hysteresis during wetting and drying cycles in soil. Soil Science of American Journal 52: 1626–1629.
- Groffman, P. M., E. A. Holland, D. D. Myrold, G. P. Robertson & X. Zou, 1999. Denitrification. In Robertson, G. P., D. C. Coleman, C. S. Bledsoe & P. Sollins (eds), Standard Methods for Long-term Ecological Research. Oxford University Press, Oxford: 272–288.
- Håkanson, L. & M. Jansson, 1983. Principles of Lake Sedimentology. Springer-Verlag, Berlin: 316 pp.
- Hedin, L. O., J. C. Von Fischer, N. E. Ostrom, B. P. Kennedy, M. G. Brown & G. P. Robertson, 1998. Thermodynamic constrains on nitrogen transformations and other biogeochemical processes at soil stream interfaces. Ecology 79: 684–703.
- Hermansson, A., J. S. K. Bäckman, B. H. Svensson & P.-E. Lindgren, 2004. Quantification of ammonia-oxidizing bacteria in limed and non-limed acidic coniferous forest soil using real-time PCR. Soil Biology and Biochemistry 36: 1935–1941.
- IPCC, 2007. Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K. B. Averyt, M.Tignor, H. L. Miller (eds)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.
- Jones, J. G., 1979. Microbial nitrate reduction in freshwater sediments. Journal of General Microbiology 115: 27–35.
- Kemp, M. J. & W. K. Dodds, 2001. Centimeter-scale patterns in dissolved oxygen and nitrification rates in a prairie

Deringer

stream. Journal of the North American Benthological Society 20: 347–357.

- Kemp, M. J. & W. K. Dodds, 2002. Comparisons of nitrification and denitrification in prairie and agriculturally influenced streams. Ecological Applications 12: 998–1009.
- Kern, J., A. Darwich, K. Furch & W. J. Junk, 1996. Seasonal denitrification in flooded and exposed sediments from the Amazon floodplain at Lago Camaleao. Microbial Ecology 32: 47–57.
- Kieft, T. L., E. Soroker & M. K. Firestone, 1987. Microbial biomass response to a rapid increase in water potential when dry soil is wetted. Soil Biology and Biochemistry 19: 119–126.
- Konhauser, K., 2007. Introduction to Geomicrobiology. Blackwell Science Ltd., Oxford: 440.
- Mitchell, A. & D. S. Baldwin, 1998. Effects of desiccation/ oxidation on the potential for bacterial mediated P release from sediments. Limnology and Oceanography 43: 481–487.
- Mitchell, A. & D. S. Baldwin, 1999. The effects of sediment desiccation on the potential for nitrification, denitrification, and methanogenesis in an Australian reservoir. Hydrobiologia 392: 3–11.
- Modini, C., M. Contin, L. Leita & M. De Nobili, 2002. Response of microbial biomass to air-drying and rewetting in soils and compost. Geoderma 105: 111–124.
- Mulholland, P. J., A. M. Helton, G. C. Poole, R. O. Hall Jr., S. K. Hamilton, B. J. Peterson, J. L. Tank, L. R. Ashkenas, L. W. Cooper, C. N. Dahm, W. K. Dodds, S. E. G. Findlay, S. V. Gregory, N. B. Grimm, S. L. Johnson, W. H. McDowell, J. L. Meyer, H. M. Valett, J. R. Webster, C. P. Arango, J. J. Beaulieu, M. J. Bernot, A. J. Burgin, C. C. Crenshaw, L. T. Johnson, B. R. Niederlehner, J. M. O'Brien, J. D. Potter, R. W. Sheibley, D. J. Sobota & S. M. Thomas, 2008. Stream denitrification across biomes and its response to anthropogenic nitrate loading. Nature 454: 202–206.
- National Assessment Synthesis Team, 2001. Climate Change Impacts on the United States: The Potential Consequences of Climate Variability and Change. US Global Change Research Program, Washington, DC: 541.
- Nielsen, L. P., P. B. Christensen, N. P. Revsbech & J. Sorensen, 1990. Denitrification and photosynthesis in stream sediment studied with microsensor and whole-core techniques. Limnology and Oceanography 35: 1135–1144.
- Pattinson, S. N., R. García-Ruiz & B. A. Whitton, 1998. Spatial and seasonal variation in denitrification in Swale-Ouse system, a river continuum. The Science of the Total Environment 210(211): 289–305.
- Peterson, B. J., W. M. Wollheim, P. J. Mullholland, J. R. Webster, J. L. Meyer, J. L. Tank, E. Marti, W. B. Bowden, H. M. Valett, A. E. Hershey, W. H. McDowell, W. K. Dodds, S. K. Hamilton, S. Gregory & D. D. Morall, 2002. Control of nitrogen export from watersheds by headwater streams. Science 292: 86–90.
- Qiu, S. & A. J. McComb, 1994. Effects of oxygen concentration on phosphorus release from reflooded, air-dried wetland sediments. Australian Journal of Marine and Freshwater Research 45: 1319–1328.
- Qiu, S. & A. J. McComb, 1995. Planktonic and microbial contributions to phosphorus release from fresh and air-

dried sediments. Marine & Freshwater Research 46: 1039–1045.

- Qiu, S. & A. J. McComb, 1996. Drying-induced stimulation of ammonium release and nitrification in reflooded lake sediment. Marine Freshwater Research 47: 531–536.
- Richardson, W. B., E. A. Strauss, L. A. Bartsch, E. M. Monroe, J. C. Cavanaugh, L. Vingum & D. M. Soballe, 2004. Denitrification in the Upper Mississippi River: rates, controls, and contribution to nitrate flux. Canadian Journal of Fisheries and Aquatic Sciences 61: 1102–1112.
- Rigobelo, E. C. & E. Nahas, 2004. Seasonal fluctuations of bacterial populations and microbial activity in soils cultivated with Eucalyptus and *Pinus*. Science and Agriculture 61: 88–93.
- Smith, M. S. & L. L. Parsons, 1985. Persistence of denitrifying enzyme activity in dried soils. Applied and Environmental Microbiology 49: 316–320.
- Solorzano, L., 1969. Determination of ammonia in natural waters by the phenolhypochlorite method. Limnology and Oceanography 14: 799–801.
- Sophocleuous, M., 2000. From safe yield to sustainable development of water resources—the Kansas experience. Journal of Hydrology 235: 27–43.
- Sorensen, J., 1978. Capacity for denitrification and reduction of nitrate to ammonia in coastal marine sediment. Applied and Environmental Microbiology 35: 301–305.
- Starry, O. S., H. M. Valett & M. E. Schreiber, 2005. Nitrification rates in a headwater stream: influences of seasonal variation in C and N supply. Journal of the North American Benthological Society 24: 753–768.
- Strauss, E. A. & W. K. Dodds, 1997. Influence of protozoa and nutrient availability on nitrification rates in subsurface sediments. Microbial Ecology 34: 155–165.
- Strauss, E. A. & G. A. Lamberti, 2000. Regulation of nitrification in aquatic sediments by organic carbon. Limnology and Oceanography 45: 1854–1859.

- Strauss, E. A., N. L. Mitchell & G. A. Lamberti, 2002. Factors regulating nitrification in aquatic sediments: effects of organic carbon, nitrogen availability, and pH. Canadian Journal of Fisheries and Aquatic Sciences 59: 554–563.
- Strauss, E. A., W. B. Richardson, L. A. Bartsch, J. C. Cavanaugh, D. A. Brusewitz, H. Imker, J. A. Heinz & D. M. Soballe, 2004. Nitrification in the Upper Mississippi River: patterns, controls, and contribution to the nitrate budget. Journal of the North American Benthological Society 23: 1–14.
- Strauss, E. A., W. B. Richardson, J. C. Cavanaugh, L. A. Bartsch, R. M. Kreiling & A. J. Standorf, 2006. Variability and regulation of denitrification in an Upper Mississippi River backwater. Journal of the North American Benthological Society 25: 596–606.
- Tiedje, J. M., 1988. Ecology of denitrification and dissimilatory nitrate reduction to ammonium. In Zehnder, A. J. B. (ed.), Biology of Anaerobic Microorganisms. John Wiley and Sons, New York, NY: 179–244.
- Triska, F. J., J. H. Duff & R. J. Avanzino, 1990. Influence of exchange flow between the channel and hyporheic zone on nitrate production in a small mountain stream. Canadian Journal of Fisheries and Aquatic Sciences 47: 2099–2111.
- Wall, L. G., J. L. Tank, T. V. Royer & M. J. Bernot, 2005. Spatial and temporal variability in sediment denitrification within an agriculturally influenced reservoir. Biogeochemistry 76: 85–111.
- Zaman, M. & S. X. Chang, 2004. Substrate type, temperature, and moisture content affect gross N mineralization and nitrification rates in agroforestry systems. Biology and Fertility of Soils 39: 269–279.